

## On distribution of radiogenic argon in minerals

BRANDT, S.B., RASSKAZOV, S.V., BRANDT, I.S. and IVANOV, A.V. Institute of the Earth's Crust SB RAS, Irkutsk, Russia.

In conventional K-Ar geochronology, bulk concentrations of K and Ar are required. Recently, interest has arisen for knowledge on distribution in mineral crystals of Ar concentration, which usually obtained by  $^{40}\text{Ar}/^{39}\text{Ar}$  modification of the K-Ar method with a laser extraction technique. We suggest a procedure of obtaining the Ar distribution by means of a macroscopic experiment. In the case of conventional K-Ar system, K is postulated to be distributed homogeneously.

We consider an aggregate of mineral grains formed  $t$  years ago and subjected to metamorphic Ar losses  $t_1$  years ago. At present, the bulk concentration of radiogenic Ar consists of two parts: one remaining just after  $t_1$  (referred to as "inherited" Ar) and another one accumulated due to K-decay in the time interval  $t - t_1$  (referred to as "normal" Ar). Due to solutions of Fick's diffusion equations the distribution of inherited Ar in the mineral is presented by a half-sinusoid, whereas the distribution of normal Ar by a rectangle. Fractions of the inherited and normal Ar in a grain with height  $h$  can be determined by a heating experiment at a constant temperature. Three portions of  $^{40}\text{Ar}_{\text{rad}}$  ( $\text{Ar}_1$ ,  $\text{Ar}_2$ ,  $\text{Ar}_3$ ), extracted at successive time intervals, allow to establish a set of algebraic equations and to calculate the ratio ( $\alpha$ ) of the inherited to normal Ar.

$$\begin{aligned} a_1 &= (1-\alpha)\Phi(\text{Fo}_1) + \alpha(1-\exp(-\pi^2\text{Fo}_2)) \text{ and} \\ a_2 &= (1-\alpha)\Phi(\text{PFo}_1) + \alpha(1-\exp(-\pi^2\text{PFo}_2)), \text{ where} \\ a_1 &= \text{Ar}_1/(\text{Ar}_1+\text{Ar}_2+\text{Ar}_3), \quad a_2 = (\text{Ar}_1+\text{Ar}_2)/(\text{Ar}_1+\text{Ar}_2+\text{Ar}_3), \quad \text{Fo}_1 = \text{Dt}_1/h^2, \\ \text{Fo}_2 &= \text{D}(t_1 + t_2)/h^2, \quad \text{Fo}_2 = \text{PFo}_1, \quad \text{P} = (t_1 + t_2)/t_1, \text{ and} \\ \Phi(\text{Fo}_1) &= 4\sqrt{\text{Fo}_1}/\pi \end{aligned}$$